The Basement of the Andean Frontal Cordillera in the Cordón del Plata (Mendoza, Argentina): Geodynamic Evolution

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ABSTRACT. The Andean Paleozoic basement of the Cordón del Plata (Argentina) consists of two sets of rocks showing different stratigraphy, structure and metamorphism. The lower one is represented by the pre-Carboniferous (Devonian?) Vallecitos beds. These rocks have been affected by folds and associated cleavage, developed under low-grade metamorphic conditions and related to the Chanic orogenic event of the Famatinian Orogenic Cycle (Upper Devonian-Lower Carboniferous). The Vallecitos beds are pre-orogenic to the Chanic deformation event and must have been deposited in a fore-arc basin located on the active margin of Chilenia, before its collision with the passive margin of Gondwana. The upper set unconformably rests on the Vallecitos beds and consists of the Late Carboniferous El Plata Formation, and the Río Blanco Conglomerates, probably Late Carboniferous-Permian in age. These rocks have been affected by east-directed thrusts and associated folds formed under very low-grade to non-metamorphic conditions. This deformation can be related to the San Rafael phase (Gondwanan Orogen). The El Plata Formation was deposited in an extensional back-arc basin while the Río Blanco conglomerates must have been deposited in a retro-arc basin, both of them on the active margin of Gondwana. The Choiyoi Group, essentially volcanic and Permo-Triassic in age, rests unconformably on the previously described successions and was deposited in an extensional setting (pre-orogenic stage) associated with the beginning of the Andean Cycle. This cycle lead to the uplift of the Frontal Cordillera during the Cenozoic and the deposition of thick continental units at the base of the Cordón del Plata in Neogene-Quaternary times.

Keywords: Paleozoic Basement, Chanic Orogeny, Gondwanan Orogeny, Geodynamic Evolution, Frontal Cordillera, Andes, Argentina.

doi: 10.5027/andgeoV39n2-a03
1. Introduction

The Cordón del Plata is a mountain range located 50 km west of Mendoza (Argentina) and is part of the Frontal Cordillera of the Andes (Groeber, 1938) (Fig. 1).

The Frontal Cordillera contains a Paleozoic basement constituted by sedimentary, metamorphic and igneous rocks (Fig. 2), which was strongly deformed during the Famatinian and Gondwanan orogenic cycles (Ramos, 1988) and is intruded by Upper Paleozoic granitoids. An Andean cover lies unconformably over the Paleozoic basement, and is constituted by Permo-Triassic and Cenozoic sedimentary, volcanic and volcaniclastic rocks (Fig. 2), intruded by Mesozoic and Cenozoic granitoids. This cover was deformed in Cenozoic times, during the Andean Orogenic Cycle. The main Andean orogeny uplift is linked with the Nazca plate subhorizontal subduction, located between 27º and 33º30’S latitude along the Pampean flat-slab segment (Gutscher et al., 2000). The absence of recent volcanism in the Pampean flat-slab segment allows the observation of the Paleozoic, Mesozoic and Cenozoic rocks and structures.

The study area is more than 6,000 m high and contains an extensive outcrop of the Paleozoic basement of the Andes (Fig. 3), formed by two sets of rocks with different deformational features and metamorphic grades. Permo-Triassic volcanic rocks of the Choiyoi Group and Cenozoic sedimentary rocks unconformably overlie this basement (Figs. 2 and 4).

The N-S trending La Carrera fault system (Campos, 1965; Polanski, 1972), whose main faults from west to east are: Arenal, Médanos and Río Blanco (Fig. 2), has been traditionally recognized as responsible for the Andean uplift of this range, although the first development of these structures has been recently assigned to the Permian San Rafael phase of the Gondwanan Orogen (Giambiagi et al., 2011). This fault system, composed by several N-S trending thrusts, uplifted the basement rocks of the Cordón del Plata range during the Cenozoic Andean Orogen. The easternmost faults of this system (between Médanos and Río Blanco faults) affect the Neogene synorogenic deposits (Folguera et al., 2003; Casa, 2005), and show evidence of Quaternary activity (Fauqué et al., 2000; Casa, 2005).

The aim of this study is to characterize the basement rocks outcropping in the eastern slope of the Cordón del Plata and compare them with studied sequences of neighbouring areas. By analyzing the deformation history and metamorphism, we attempt to establish the tectono-stratigraphic evolution of these rocks and fit them into a coherent evolution model for the Andes during the Late Paleozoic.

2. Stratigraphy

There are four main stratigraphic units, separated by unconformities in the study area (Figs. 2 and 4):

I. The oldest unit consists of centimeter to decimeter scale, well-stratified beds of dark quartzites and laminated shales (Fig. 5A), where cross-and flaser-laminations can often be recognized. Meter thick beds of clast-supported conglomerates, including centimeter-sized clasts of quartz, plutonic, volcanic and other metamorphic rocks can occasionally be observed. Until now, no sedimentologic studies have been made on these rocks, although their stratigraphic characteristics could indicate a turbiditic origin.

The age of these rocks is not well known due to the absence of fossils. Based on lithological characteristics and stratigraphic position, this sequence is very
similar to the Alto Tupungato Formation, described by Polanski (1959) in the western part of the Cordón del Plata massif, which has been assigned to Early Carboniferous or older. However, this formation was attributed to the Late Carboniferous (López Gamundi et al., 1987), which would invalidate this possible correlation. In addition, successions with similar facies, covered unconformably by Late Carboniferous rocks, have been assigned to the Devonian in the Frontal Cordillera of the San Juan province (Gutiérrez, 1983; Azcuy et al., 1999; Heredia et al., 2002). Recently, south of the study area, Tickyj et al. (2009) have described Late Ordovician beds in the eastern side of the Frontal Cordillera (Las Lagunitas Formation); but their stratigraphic features are different (more slates and lithic sandstone predominance) from those described for the Cordón del Plata area and they could be older. For all these reasons, we provisionally named these rocks as Vallecitos beds and assigned them to the Devonian.

The absence of subaerial volcanism older than Late Carboniferous in the eastern part of the Frontal Cordillera (including the Precordillera), suggests a western origin for the volcanic clasts included in the conglomerates of the Vallecitos beds.

This succession is intruded by many deformed dykes of quartz and pegmatites of granitic/granodioritic composition, which do not affect the overlying El Plata Formation.

The Vallecitos beds have similar facies to the Devonian successions described in the central and eastern Precordillera, which are part of the Gualilán Group (Baldis and Sarudiansky, 1975) (Fig. 6). The same Devonian age has been assigned to the Villavicencio and El Codo formations, described in the western part of the Precordillera (Harrington, 1941; Guerstein et al., 1965; Sessarego, 1988; Baldis and Peralta, 1999; Amenabar and di Pasquo, 2004). In contrast to the Vallecitos series, the Gualilán Group has an eastern non-volcanic source area (González Bonorino, 1975),
FIG. 3. Photographies showing the main geological features of the east side of the northern (A) and southern (B) Cordón del Plata. 

VAS: Vallecitos beds; EPF: El Plata Formation; i: lower member; m: middle member; s: upper member; CHG: Choiyoi Group; 
GR: Cerro Arenal granite; NEO: Neogene deposits. White lines: lithostratigraphic contacts; Dotted white lines: Apparent 
lithostratigraphic boundaries in the picture; Black triangles: Gondwanan thrust; Blank triangles: Andean thrust.
which hamper the correlation between the two stratigraphic units. As in the Vallecitos beds, the Ratón Formation from the western Precordillera, included in the Angualasto Group (Early Carboniferous) (Fig. 6), contains conglomerates with volcanic clasts of western origin. Although the Angualasto Group and the Vallecitos beds have similar deformational style, these units do not seem to correlate, since the Angualasto Group shows different facies and unconformably overlies the Devonian series (El Codo Formation).

II. A thickening siliciclastic succession rests unconformably over the Vallecitos beds. It is composed of the El Plata Formation (Caminos, 1965, 1979) at the bottom part and the Río Blanco Conglomerates at the top (Fort, 1944; Caminos, 1965).

The El Plata Formation has a thickness of nearly 3,300 m in the study area, and includes three members (Caminos, 1965):

- The lower member crops out along several bands bounded by thrusts in the lower part of the eastern slope of the Cordón del Plata massif (Fig. 3A) and it can reach a thickness of 1,200 m in nearby areas. Its lower part consists of conglomerates with plutonic and metamorphic pebbles. In the upper part, the conglomerates change to alternating arkosic and quartz sandstones, interbedded with gray-green shales and some quartzitic microconglomerates. This member is very similar to the Loma de los Morteritos Formation, defined by Polanski (1959) in nearby areas, and is assigned to the Late Carboniferous (Folguera et al., 2003).

- The middle member is almost 1,500 m thick and is composed of gray medium to coarse-grained quartzitic litarenites, in meter to decimetre thick beds with thin interbedded shales. These rocks form the eastern wall of the Vallecitos and Rincon peaks and the east-southeast face of the El Plata peak (Figs. 3 and 5B).

- The upper member can exceed 4,000 m in thickness (Azcuy et al., 1999), although in the study area it does not exceed 600 m as it is covered by thrusts. This member is mainly composed of shales (Fig. 5C) with centimetre thick intercalations of fine-grained sandstones with ripple lamination towards the top. At its base, it contains a very characteristic carbonaceous shale intercalation that is intensely tectonized.
The Río Blanco Conglomerates (Fort, 1944; Caminos, 1965) crop out in the La Manga Quebrada, located in the northern part of the study area (Fig. 2). Due to their moderate thickness of 20-80 m, they have not been represented on the geological map of the figure 2. These conglomerates contain metamorphic, sedimentary and volcanic pebbles and unconformably overlie the El Plata Formation, being unconformably covered by the Tambillos Formation (Choiyoi Group) of lower Permian age (Cortés et al., 1997).

The El Plata Formation is similar in lithology and stratigraphic characteristics to the Agua Negra Formation (Polanski, 1970), which outcrops in the Frontal Cordillera to the west of San Juan province (Fig. 6). Moreover, the lower part of the San Ignacio Formation (Carboniferous-Permian), which overlies unconformably the Agua Negra Formation (Heredia et al., 2002; Busquets et al., 2005), is similar to the Río Blanco Conglomerates (Fig. 6). Both formations, San Ignacio and Río Blanco, could represent an alluvial fan system in which volcanic levels are interbedded.

### III. The Paleozoic basement is unconformably covered by the Choiyoi Group volcanics (Figs. 4, 7 and 9C) of Permian-Triassic age (Groebler, 1946; Rolleri and Criado, 1969). In the study area, this group is composed of acidic to intermediate volcanic and volcaniclastic rocks as well as plutonic or subvolcanic intrusions with the same composition (Cortés et al., 1997). The upper part is composed of sedimentary rocks (only in the Precordillera, Fig. 2), that are mainly constituted by clastic continental deposits with some interbedded volcanic rocks.

The Paleozoic and Early Mesozoic rocks are intruded by Triassic granitoids that produced a thermal metamorphism with amphibole and biotite hornfels. The outcrops of these granitoids show an elongated N-S trend (Fig. 2 and 4), running parallel to the Andean structures. Caminos et al. (1979) dated Cerro Arenal and Cerro Médanos (Precordillera) batholiths in the study area as Middle Triassic (K-Ar ages of 237±10 and 244±10 Ma respectively).

### IV. Finally, in the foothills of the Cordón del Plata massif, Cenozoic (Neogene) conglomerates, sandstones and shales of continental origin rest unconformably on the rocks described above. They are part of the alluvial fans that drain the relief of the Frontal Cordillera (Figs. 2 and 3A) and they represent the synorogenic deposits of the Andean Cycle.

Freije et al. (1999) described basaltic pillow lavas, up to 5 m thick, interbedded in the sediments of the El Plata Formation. The sedimentary environment of the El Plata Formation ranges from shallow marine platform facies in the lower member to fluvial and deltaic in the middle and upper ones respectively (Folguera et al., 2003).
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### FIG. 6. Stratigraphic correlations panel.

- **Fm**: formation
- **Gr**: group
- **Or**: orogeny
- **Cg**: conglomerates

*In Italics*: the equivalent lithostratigraphic units of nearby areas.

### FIG. 7. Geological cross section.

- **VAS**: Vallecitos beds
- **EPF**: El Plata Formation
- **i**: lower member
- **m**: middle member
- **s**: upper member
- **CHG**: Choiyoi Group
- **GR**: Cerrito Arenal granite
- **NEO**: Neogene deposits

Horizontal and vertical scale is the same.
3. Structure and Metamorphism

The pre-Andean basement of the Cordón del Plata massif is affected by Paleozoic to Cenozoic structures, developed during four deformation episodes. The oldest one (I) took place in pre- to Early Carboniferous times under low-grade metamorphic conditions and can be linked to the Famatinian Orogenic Cycle. The second episode (II), Late Carboniferous-Permian in age, was developed in very low grade to non-metamorphic conditions during the Gondwanan Cycle. The last two episodes (III and IV) belong to the Andean Cycle and took place without metamorphism.

3.1. Pre-Late Carboniferous Famatinian structures

This deformation episode only affects Vallecitos beds and is characterized by centimeter to decimeter scale close to tight subvertical folds. These structures are folding an S1 slaty cleavage (Figs. 8A and B), which forms an angle lower than 10° with bedding (Fig. 8A). In the shaly layers, S1 is defined by shape preferred orientation of chlorite and muscovite porphyroblasts and quartz crystals. In the quartzites, S1 is only recognized by the orientation of some phyllosilicates and opaque minerals (Fig. 8C), and can be classified as a spaced disjunctive cleavage. In the microlithons, quartz is accompanied by plagioclase and K-feldspar. The slaty cleavage (S1) is usually crenulated (Figs. 8A and C) and sometimes develops a crenulation cleavage (S2). At the microscopic scale, S2 is a rough, anastomosing and discontinuous crenulation cleavage (Fig. 8C), associated to the folds (Fig. 8A). Both foliations, S1 and S2, are sometimes partially obliterated by the growth of biotite and amphibole associated to the thermal Triassic metamorphism, related to magmatic activity.

3.2. Late Carboniferous-Permian Gondwanan structures

These structures are represented by westward dipping and east-directed thrusts (Fig. 4 and 7) with associated subhorizontal N-S meter sized folds. Some of the folds have an associated rough cleavage in shaly beds. The thrust detachment is located in the basal member of the El Plata Formation, in the easternmost sector of the El Plata range, but in the western part it develops in the basal sandstones of the middle member (Figs. 3B, 4 and 7). Another detachment is located at the base of the upper slate member of the El Plata Formation (Fig. 5C, 7 and 9B). Some of these thrusts produce a kilometer scale shortening, which is greater than 5 km in the La Jaula thrust. An out-of-sequence thrust, El Salto Thrust, cuts a previous one producing the
folding of the La Jaula thrust (Figs. 3, 4, 7 and 9A) and is covered unconformably by the Choiyoi Group, allowing the dating of these structures as pre-Andean (Fig. 4).

3.3. Andean extensional structures

From the Late Permian to the Middle Triassic this part of the Frontal Cordillera was affected by extensional tectonics (Mpodozis and Ramos, 1989; Uliana et al., 1989; Llambias et al., 1993) that started the Andean Cycle (Heredia et al., 2002). The Permo-Triassic normal faults show a N-S preferred trend, with some segments oriented NE-SW to NW-SE (Figs. 2, 4 and 10). These faults are contemporaneous with the deposition of the Choiyoi Group, and produced significant variations in its thickness and lithological characteristics (Heredia et al., 2002). The elongated geometry of the Triassic granitoids (Figs. 2 and 4), with the same trend of the normal faults, indicates that their intrusion could be structurally controlled by these faults.

The Vallecitos-La Hoyada normal fault must have developed during this period, as the Choiyoi Group does not exist to the south and east of this structure (Fig. 4). The change in the trend of the fault, from N-S to E-W (Figs. 4 and 10), could be controlled by the presence of pre-Andean structures with the same directions. This fault is very steep in the segment with E-W trend (Vallecitos Fault, Fig. 10), allowing to interpret it as a transfer zone of the Permo-Triassic extensional system, which mainly trends N-S in this area of the Frontal Cordillera (Fig 10).

The Río Blanco, Cerro Médanos and Cerro Arenal faults (Figs. 2, 3A and 7), located at the eastern border of the Cordón del Plata massif, could also be considered as related with the Permo-Triassic extension (Fig. 10), although they were reactivated and inverted during later Cenozoic compression. In the northern part of the Río Blanco fault, coinciding with the southern outcrops of the Precordillera, the thickness of the Choiyoi Group is similar in both the eastern and western fault walls. However, the Choiyoi Group in the southern part of the fault is very thin, and is only represented by the lower volcanic part, or is even entirely absent in the eastern hanging wall (Fig. 7). This coincidence could be explained by the existence of an E-W trending transfer fault (Fig. 10), which is now covered by Cenozoic rocks (Potrerillos...
fault in this paper, Fig. 2). Finally, in the hanging wall of the N-S trending Cerro Medanos fault, the thickness of the Choiyoi Group is around 2,000 m, the common thickness in the Frontal Cordillera, where it can exceed 4,000 m (Heredia et al., 2002).

3.4. Andean compressive structures

In the Cenozoic, the Cerro Arenal, Médanos and Río Blanco faults produced the main uplift of the Cordón del Plata massif juxtaposing the pre-Andean basement and the Neogene rocks (Cerro Médanos fault, Figs. 2, 3A and 7). The Andean compressive faults are usually subvertical and merge in a common detachment dipping to the west (Folguera et al., 2003, Giambiagi et al., 2010) (Fig. 7). Displacements of these faults range from hundreds of meters to kilometers (Fig. 7). As shown in the geological map and cross-sections (Figs. 2, 4 and 7), the Andean faults reactivated some Gondwanan thrusts and Mesozoic extensional structures. In this way, the La Hoyada-Vallecitos normal fault could also have been active during the Andean Orogeny as a lateral ramp, as indicated the absence of back-thrusts in the southern block. The Cenozoic deformation could also be responsible for the Potrerillos Fault reactivation as a lateral ramp of an Andean thrust. Such reactivation could cause the steep plunge of the folds axes in the Precordillera and its abrupt southern termination (Fig. 2).

4. Proposal of Geodynamic Evolution during the Late Paleozoic

In recent years, several models on the Late Paleozoic geodynamic evolution of the southern Central Andes have been published (Ramos et al., 2000; Davis et al., 2000; Ramos, 2004; Massonne and Calderón, 2008). In these models, the geodynamic evolution is linked to the accretion of the Chilenia terrane (Fig. 1) to the Gondwana margin, formed by the Cuyania terrane (Fig. 1) previously accreted during Silurian-Early Carboniferous times (Famatinian Orogenic Cycle). The subsequent development of a subduction in the western margin of the Chilenia terrane was mainly developed during de Gondwanan
Orogenic Cycle in Late Carboniferous-Early Permian times (Rebolledo and Charrier, 1994).

New details of the geodynamic evolution of the southern Central Argentine Andes during the Late Paleozoic can be obtained from our study; based on the presence in the Cordón del Plata of three rock sets with different structural and metamorphic history, separated by two major unconformities (Figs. 2 and 6). Our data was completed and compared with that obtained by different authors in surrounding areas of the Cordillera Frontal and also in the Precordillera.

The main structures affecting the rocks of the Vallecitos beds in the Cordón del Plata must have been generated during the Famatinian Orogenic Cycle. The possible Devonian age of these rocks and their location in the eastern part of the Frontal Cordillera, suggest that their structures are related to the Chánic Orogeny of that cycle, produced by the accretion of Chilenia to the Gondwana margin, formed by the Cuyania terrane (Fig. 1), previously accreted in Late Devonian times (Ramos et al., 1984). This collision left remnants of Late Proterozoic-Silurian rocks with ophiolite affinities on the western side of the Precordillera, (Davis et al., 1999) that marked the separation (suture) between these two terranes (Fig. 1). If this interpretation is correct, Vallecitos beds should correspond to sediments deposited on the continental margin of Chilenia (Fig. 11A) while the coeval series in the Precordillera (Fig. 5) were deposited on the continental margin of Gondwana (Fig. 11A). The volcanic clasts in the Vallecitos beds point to the existence of an active volcanic area in the Chilenia terrain. This volcanism would imply a west-dipping subduction, in which the oceanic crust that separated the two continents was completely consumed under Chilenia (Davis et al., 1999; Gerbi et al., 2002; Giambiagi et al., 2009), although the separation between both continents should not have been very wide (Rapela et al., 1998; Busquets et al., 2005). In addition, Tickyj (2011) describes Lower Devonian calc-alkaline granitoids (Fig. 11A) in the eastern part of the Frontal Cordillera (Cordón del Carrizalito), just south of the study area, which have been interpreted as belonging to a magmatic arc. In this context, the rocks of Vallecitos beds were deposited in the active margin of Chilenia, probably in a fore-arc basin associated with the subduction (Fig. 11A). It has not been possible to deduce the vergence of the Chánic structures in the study area. West-verging structures of this age (Fig. 11A) have been described in northern parts of the Frontal Cordillera (Heredia et al., 2002) and also further south (Tickyj et al., 2009). On the other hand the presence in the Frontal Cordillera, further to the south and east, of high-grade metamorphic rocks (Bjerg et al., 1990) and granitoids (Caminos et al., 1979; Tickyj, 2011) of Early Devonian to Early Carboniferous age, point to the presence of the Chànic Orogenic hinterland in the eastern Frontal Cordillera. In contrast, the Chánic structures in the Central and Western Precordillera show a general vergence towards the east (Davis et al., 1999; Álvarez Marrón et al., 2006; Alonso et al., 2008) and the metamorphic grade increases westward (Von Gosen, 1992), as it might be expected from its location on the Gondwana passive margin (Alonso et al., 2008), prior to the collision of Chilenia (Fig. 11A).

On the other hand, no Chánic synorogenic rocks have been found in the study area, as in the rest of the Frontal Cordillera. However in the Precordillera, these rocks are well represented by the Early Carboniferous Angualasto Group (Limarino and Cesari, 1992, Limarino et al., 2006) (Fig. 6). The presence in this stratigraphic unit of conglomerates with volcanic and plutonic clasts of western provenance and Early Carboniferous age, represents the erosion of the magmatic arc from Chilenia during the Chánic Orogeny (Fig. 11B). The Chánic plutonic activity must be linked with the subduction in pre-collisional times and with the crustal thickening that occurred during the orogenic process, in syn-collisional times. The presence of low-grade regional metamorphism, scarce plutonism and westward verging structures in nearby areas, supports the conclusion that this sector of the Frontal Cordillera was located in the hinterland of the western segment of the Chánic Cordillera (Fig. 11B).

After the Gondwana-Chilenia collision, a new subduction zone developed on the western margin of accreted Chilenia (Ramos, 1988; Rebolledo and Charrier, 1994), beginning the Gondwanan Orogenic Cycle (Fig. 11C). In the Late Carboniferous, the sedimentation of the El Plata Formation took place in a back-arc extensional basin (Fig. 11C), representing the first Gondwanan pre-orogenic sequence (Mpodozis and Ramos, 1989; Fernández Seveso et al., 1993; Astini, 1996; Azcuy et al., 1999; Charrier et al., 2007). The Carboniferous rocks of the study area are part of the Río Blanco-Calingasta-Uspallata basin (Amos and Rollieri, 1965). The Precordillera (Amos and
FIG. 11. Geodynamic evolution sketch of the Andes between 30° and 33°S latitude, during Late Paleozoic times. Figures are not to scale. **Blank Arrows:** main areas of sediment supply. **Black arrows:** other areas of sediment supply; **Fm:** formation; **Gr:** group; **Cg:** conglomerates.
Rolieri, 1965) (Fig. 11C) developed during most of the Late Carboniferous and appears to have been a horst-like topographic high that was inherited from the Chanic Cordillera and separated the Río Blanco basin from the Paganzo basin farther east (Azcuy and Morelli, 1970; Salfitty and Gorustovich, 1983; López Gamundi et al., 1987; Azcuy, 1985; Fernández Seveso et al., 1993; Azcuy et al., 1999; Limarino y Spalletti, 2006; Limarino et al., 2006). The change in lithologic characteristics described earlier, of time-stratigraphic equivalent rocks (i.e., Agua Negra, El Plata, and Jarillal Formations in Fig. 6) reflects the relative influence of this Protoprecordillera high, that appears to have become progressively less prominent to the south and was absent south of the study area. These Gondwanan pre-orogenic rocks are mainly sourced from the East, in the present-day Sierras Pampeanas (Limarino, 1987; Heredia et al., 2002). However, where the Protoprecordillera was present, it provided a local source of sediment. On the far west, in Río Blanco basin, a western source of sediment (from the volcanic arc) can be recognized, but it is rare because, the arc was not well developed at this time (Fig. 11C). Río Blanco basin probably began as an extensional back-arc feature that was initially filled largely with marine deposits. The pre-orogenic extensional character of this basin is compatible with the moderate calc-alkaline volcanism found in the Frontal Cordillera and Precordillera (Freije et al., 1999; Koukharsky et al., 2009) and the alkaline volcanism (Koukharsky et al., 2002) in eastern areas, away from the Gondwanic subduction zone (Paganzo Basin) (Fig. 11C). In the later phases of the Gondwanan Cycle (San Rafael orogenic phase of Late Carboniferous-Permian times) contractional deformation, with east-directed thrusts and folds, produced a retro-arc foreland basin that was filled with synorogenic continental deposits (Río Blanco Conglomerate in the Cordón del Plata area and San Ignacio Formation farther north). These synorogenic deposits (Fig. 6) contain clasts of volcanic origin and, only in the San Ignacio Formation, volcanic and volcano-sedimentary rocks sourced from the active volcanic arc. Moreover, the volcanic intercalations in these successions indicate that the volcanic arc remained active during the entire orogenic process (Fig. 11D). Except in the Paganzo Basin, the source area of the Gondwanan synorogenic deposits was located to the west, forming part of a retro-arc foreland basin that shows isolated outcrops at present (Busquets et al., 2005). The Protoprecordillera disappeared at this time (Fig. 11D), and the El Salto Formation, the Pircas Conglomerates and other equivalent synorogenic series (Fig. 6) were deposited in the Precordillera area (Fig. 11D).

Considering the deformation style and the absence of metamorphism and plutonism, the Frontal Cordillera can be interpreted as part of the foreland of the Gondwanan Orogen, which was far away from the subduction zone (Fig. 11D). Finally, the Choiyoi Group rests unconformably over all the rocks described above and represents the Permo-Triassic extensional episode (Ramos, 1988; Uliana et al., 1989; Llambias y Sato, 1990), which started the Andean Orogenic Cycle (Heredia et al., 2002).

Acknowledgements

We are grateful to reviewers F. Hogn, J. Kley and C. Schmidt, whose comments have improved the original manuscript and also to P.G. Masyby for the English review. This work has been supported by CGL2006-12415-C03/BTE and CGL2009-13706-C03 projects (Spanish I+D+i Plan).

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