

Comment on 'Geochronology (Ar-Ar, K-Ar and He-exposure ages) of Cenozoic magmatic rocks from northern Chile (18-22°S): implications for magmatism and tectonic evolution of the central Andes' of Wörner *et al.* (2000)

Marcelo García

Servicio Nacional de Geología y Minería,
Av. Sta. María 0104, Providencia, Santiago, Chile
mgarcia@sernageomin.cl

Gérard Hérail

Institut de Recherche pour le Développement,
209-213 Rue La Fayette, 75010, Paris, France

Wörner *et al.* (2000) presented new and important geochronologic data on the Cenozoic magmatic rocks of northern Chile, and proposed magmatic and tectonic interpretations for them. Here we comment on one of their interpretations about the Miocene tectonic evolution of the Precordillera of Arica (Pampa Oxaya, ca. 18°30'S), a region referred by them as 'Oxaya Block'. According to their geochronological data and geomorphological observations, the area, which mainly consists of Lower Miocene ignimbrites of the Oxaya Formation, is interpreted as a large gravitationally collapsed rotational block. They interpret the collapse to have been produced by gravitational instability due to steepening of the western Andean slope during uplift of the Altiplano in an overall compressive regime.

PROBLEMS WITH THE GRAVITATIONAL COLLAPSE MODEL

We argue that the Pampa Oxaya can not be considered as a gravitationally collapsed block because of (in order of importance):

Geometry and morphology

In the field, there is no evidence for a basal detachment, lateral boundaries, or even for an amphitheatre of a possible collapse structure.

Characteristically, source areas of collapsed blocks present a concave down-slope geometry (*e.g.*, Siebert, 1984; Philip and Ritz, 1999). According to figure 5 of Wörner *et al.*, that is not the case for the proposed amphitheatre of the Oxaya Block which shows a slightly convex down-slope shape. Furthermore, they suggest that the Ausipar Fault represents the main basal detachment at the front of the supposed collapse, which is shown with a concave-upward shape, and breaking up to the surface (Fig. 5). However, where the Ausipar Fault crops out well (Lluta Valley), it is a concave-downward fault, which dips 40-50°E in the floor of the valley and shallows to a subhorizontal thrust up to a tip line at higher levels on the valley walls (García *et al.*, 1999). Above this tip line the Oxaya Formation is only folded (Figs. 1 and 2). On the other hand, the Ausipar Fault is not connected to lateral boundaries proposed for the Oxaya Block (Fig. 5 of Wörner *et al.*, 2000), and the topography and stratigraphy are the same on both sides of these inferred lateral boundaries.

Internal structure

The surface morphology of the Pampa Oxaya is very regular. The Oxaya ignimbrites are coherent; they do not show a chaotic or disturbed structure with 'hummock' blocks and fragmented matrix, typical of gravitational slides and due to fragmentation

during transport (Siebert, 1984; Ui *et al.*, 1986). In particular, such fragmentation structures are not observed at all at the distal areas of the supposed collapse where they would be expected to be best developed.

Slope

The authors argue that the rocks at Pampa Oxaya collapsed where the Andean flank 'is by far the steepest'. However, the slope between the front (Ausipar Fault) and source (Chapiquiña-Belén ridge) of the hypothetical collapse is similar to the one observed farther north and directly south of Pampa Oxaya (from ca. 2,000 m to 4,500-5,000 m altitude of 50 km horizontal distance). In addition, the giant escarpment in the source area, as shown in the cross section of figure 5, is mainly an artifact due to its vertical exaggeration (4:1).

Volume

The volume of the Oxaya Block is about 500 km³ (estimated from Fig. 5 of Wörner *et al.*, 2000). The largest gravitational collapses known in the world, in a continental environment, have maximum volumes of 50-100 km³ (Ibetsberger, 1996; Dade and Huppert, 1998; Philip and Ritz, 1999), which are 5 to 10 times smaller than the volume estimated for the Oxaya Block.

THE PAMPA OXAYA AS FORMED BY FOLDING

The Oxaya Anticline

We agree with Wörner *et al.* (2000) that lithospheric loading cannot explain the eastward tilting of the Pampa Oxaya. However, this tilting is not necessarily produced by gravitational rotation. Alternatively, our field mapping (1:50,000 scale) indicates that the considered segment of the Precordillera presents all the characteristics of an asymmetric major gentle fold: the Oxaya Anticline. This west-vergent compressive tectonic structure was recognised by Salas *et al.* (1966), Muñoz and Charrier (1996), García *et al.* (1999) and Rochat *et al.* (1999). The fold affects the Oxaya Formation, which to the west conformably overlies the Azapa Formation, and to the east unconformably overlies a more deformed Mesozoic sequence and Cretaceous-Paleogene intrusions (Salas *et al.*, 1966) (Fig. 1). Along its strike (N10-25W), the Oxaya

Anticline can be followed for 50 km, with a half-length wave of 25 to 30 km and a maximum amplitude of 1,000 m (Fig. 1). In the hinge, buckling of the strongly competent Oxaya ignimbrites has caused normal faulting and 'hinge grabens'. These normal faults are parallel to the fold axis and show a negligible superficial displacement (up to 80 m). To the east, the Oxaya Anticline is cut by a Miocene west-vergent fold-and-thrust system (Muñoz and Charrier, 1996). The gravels of the Huaylas Formation (Salas *et al.*, 1966), overlying the eastern limb of the Oxaya Anticline, were deposited immediately after folding, indicating a Late Miocene age for the structure (García *et al.*, 1999).

The Ausipar Fault

We interpret the Oxaya Anticline as having formed by propagation at depth of the blind Ausipar Fault (Figs. 1 and 2). This thrust does not cut the upper Oxaya Formation, but its projection at the surface is well defined in the field, with a constant N20W strike. The fault projection is parallel to, and has the same extension as, the Oxaya Anticline. As described above, the Ausipar Fault is well exposed in the Lluta valley (Tiñare) (Fig. 2). At the floor of the valley, it juxtaposes the Mesozoic-Paleocene substratum with the lower Oxaya Formation. Higher up it dies into a tip line at ~1,500 m altitude. This geometry defines a fault-propagation fold involving the basement (Nair and Suppe, 1994; Mitra and Mount, 1998). The negligible horizontal shortening (60-80 m), the relatively important vertical displacement (ca. 1,000 m), and the shape and magnitude of the uplift surface in section, produced by the folding, implies that the Ausipar Fault projects downward as an east-dipping high-angle reverse fault (Fig. 1).

Comparison with the Lluta Avalanche

Gravitational instability due to uplift of the Oxaya Anticline produced, on its western flank (Lluta valley region), the Lluta Avalanche (Naranjo, 1993; Uhlig, 1999). In contrast to the hypothetical Oxaya Block, this landslide shows a clear concave down-slope amphitheatre, well-defined lateral boundaries, a basal detachment, and highly chaotic and disturbed deposits. Estimates on the volume of the Lluta Avalanche range between 50 km³ (Uhlig, 1999) and 100 km³ (Naranjo, 1993).

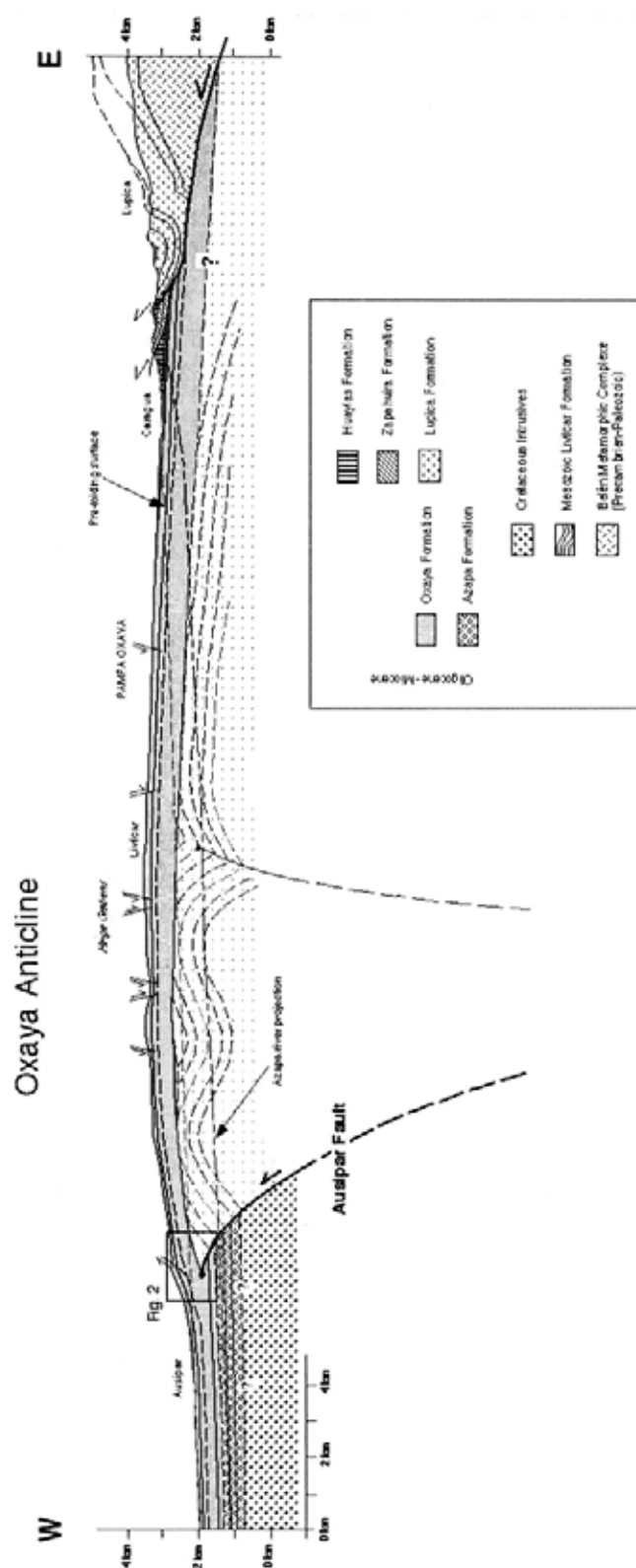


FIG. 1. Cross section of the Oxaya Anticline. No vertical exaggeration. Location is as the cross section of figure 5 in Wörner *et al.* (2000). The box shows the localisation of figure 2 projected into this cross section.

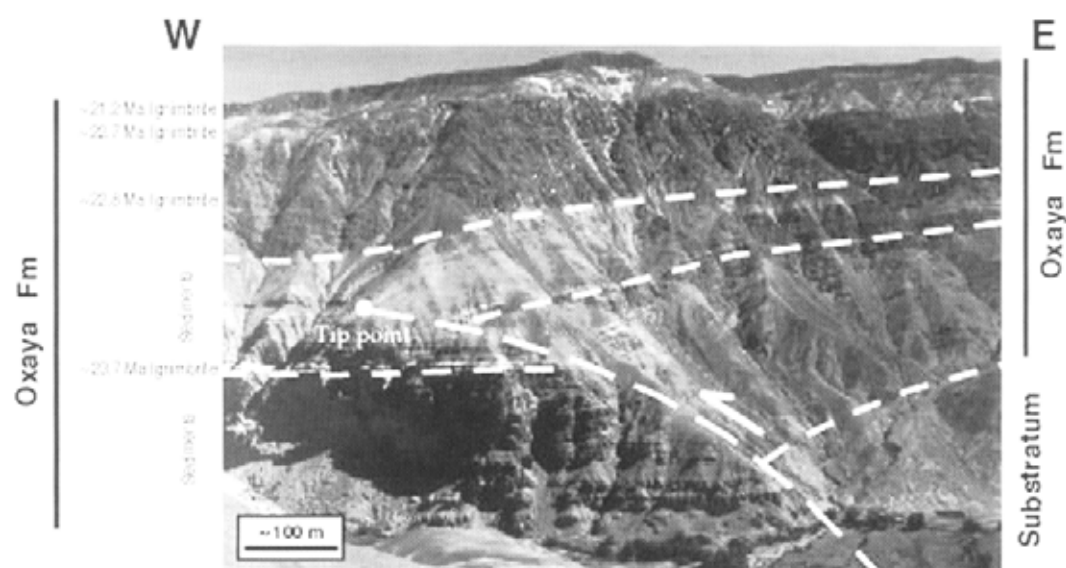


FIG. 2. The Ausipar Fault in the northern side of the Lluta valley (Tiñare). View to the North. The thinning to the east of the Oxaya Formation is apparent due to the perspective of the photography.

REFERENCES

- Dade, B.; Huppert, H. 1998. Long-runout rockfalls. *Geology*, Vol. 26, p. 803-806.
- García, M.; Hérail, G.; Charrier, R. 1999. Age and structure of the Oxaya Anticline, a major feature of the Miocene compressive structures of northernmost Chile. In *International Symposium on Andean Geodynamics*, No. 4, *Extended Abstracts Volume*, p. 249-252. Göttingen.
- Ibetsberger, H.J. 1996. The Tsergo Ri Landslide: an uncommon area of high morphological activity in the Langthang valley, Nepal. *Tectonophysics*, Vol. 260, p. 85-93.
- Mitra, S.; Mount, V. 1998. Foreland basement-involved structures. *American Association of Petroleum Geologists, Bulletin*, Vol. 82, p. 70-109.
- Muñoz, N.; Charrier, R. 1996. Uplift of the western border of the Altiplano on a west-vergent thrust system, Northern Chile. *Journal of South American Earth Sciences*, Vol. 9, p. 171-181.
- Naranjo, J.A. 1993. Hallazgo de una gigantesca avalancha de detritos del Cenozoico superior en Oxaya, Región de Tarapacá. *Segundo Simposio Internacional de Estudios Altiplánicos*, p. 47-52. Arica, Chile.
- Narr, W.; Suppe, J.; 1994. Kinematics of basement-involved compressive structures. *American Journal of Science*, Vol. 294, p. 802-860.
- Philip, H.; Ritz, J-F. 1999. Gigantic paleolandslide associated with active faulting along the Bogd Fault (Goby-Altay, Mongolia). *Geology*, Vol. 27, p. 211-214.
- Rochat, P.; Hérail, G.; Baby, P.; Mascle, G. 1999. Bilan crustal et contrôle de la dynamique sédimentaire sur les mécanismes de formation de l'Altiplano. *Comptes Rendus de l'Académie des Sciences*, Vol. 328, p. 189-195. Paris.
- Salas, R.; Kast, R.; Montecinos, F.; Salas, I. 1966. Geología y recursos minerales del Departamento de Arica, Provincia de Tarapacá. *Instituto de Investigaciones Geológicas, Boletín*, No. 21, 114 p.
- Siebert, L. 1984. Large volcanic debris avalanches: characteristics of source areas, deposits, and associated eruptions. *Journal of Volcanology and Geothermal Research*, Vol. 22, p. 163-197.
- Uhlig, D. 1999. Die Westabdachung der Zentrallanden in den Provinzen Arica und Parinacota, Nordchile: Landschaftsentwicklung und Geologie. Dissertation *University of Stuttgart*, Profil, Vol. 17, p. 167-244.
- Ui, T.; Kawachi, S.; Neall, V.; 1986. Fragmentation of debris avalanche material during flowage-evidence from the Pungarehu Formation, Mount Egmont, New Zealand. *Journal of Volcanology and Geothermal Research*, Vol. 27, p. 255-264.
- Wörner, G.; Hammerschmidt, K.; Henjes-Kunst, F.; Lezaun, J.; Wilke, H. 2000. Geochronology (Ar-Ar, K-Ar and He-exposure ages) of Cenozoic magmatic rocks from northern Chile (18-22°S): implications for magmatism and tectonic evolution of the central Andes. *Revista Geológica de Chile*, Vol. 27, p. 205-240.

Reply to the Comment by M. García and G. Hérail on
'Geochronology (Ar-Ar, K-Ar and He-exposure ages) of
Cenozoic magmatic rocks from northern Chile (18-22°S):
implications for magmatism and tectonic evolution of
the central Andes' by Wörner *et al.* (2000)

Gerhard Wörner

Abteilung Geochemie, GZG, Universität Göttingen,
Goldschmidtstr. 1, 37077 Göttingen, Germany
gwoerne@gwdg.de

Hartmut Seyfried

Institut für Geologie und Paläontologie, Universität Stuttgart,
Herdweg 51, 70174 Stuttgart, Germany

Process and timing of uplift of the Altiplano of the Central Andes is a first-order geological problem. It has attracted the attention from different disciplines such as geophysics, paleobotany, sedimentology, structural geology and geochemistry (*e.g.*, Isacks, 1988; Gregory-Wodzicki, 2000; Gaupet *et al.*, 1999; Allmendinger *et al.*, 1997; Wörner *et al.*, in press). García and Hérail commented on a paper by Wörner *et al.* (2000a), which is mostly concerned with the presentation and interpretation of geochronological data on volcanic and a few intrusive rocks in northernmost Chile. One, albeit important aspect of the interpretation of the new geochronological data relates to the tectonic evolution of the Western Andean Escarpment near Arica during the Miocene, especially the Pampa de Oxaya (*ca.* 18°30'S), which we interpreted as a large gravitationally rotated block.

AGREEMENTS

We would first like to identify those facts and interpretations where we think we agree with García and Hérail:

1- the Oxaya structure is a highly unusual morphological feature which is observed (at least on land) only in the Arica Bend where, as we have argued, the morphological gradient from coast to the crest of the Western Cordillera is greater than anywhere else in northern Chile. The surface of the

Oxaya Block has been rotated towards the east, thus creating the unusual staircase morphology of the western slope of the Andes (Fig. 1b). The angle of the inclined, eastward tilting surface, is low, only about 2 or 3 degrees.

2- the stratigraphic framework of sedimentary, pyroclastic and volcanic rocks comprises:

a- The Azapa Formation which includes the Oxaya Ignimbrites of about 25 to 19 Ma age

b- the younger Huaylas Formation overlies the Oxaya Block. Accommodation space and sedimentary infill was provided in response to the tectonic movements of the Oxaya Block and concomitant uplift of the Western Cordillera.

c- we agree that the Lupica Formation is lithologically different from the Azapa and Oxaya Formations (see figure 1 of García and Hérail's comment). This stratigraphic sequence is also consistent with that of Salas *et al.* (1966) and Seyfried *et al.* (2000). The Lupica Formation underlies the Azapa and Oxaya Formations and thus is older (possibly early Tertiary).
d- displacement and rotation of the Oxaya Block (regardless of its cause) must have occurred in the Late Miocene, *i.e.*, during the time-span between 12 Ma and 10.5 Ma (Wörner *et al.*, 2000).

3- along steep mountain fronts compressional and extensional movements may develop simultaneously indeed (Bailey, 1998).

DISAGREEMENTS

The main point of disagreement with García and Hérail concerns the question whether the rotated Oxaya Block is the result of a ramp-and-thrust structure or was caused by gravitational movements. In this context, we will discuss the remaining contentious points, some of which may turn out to be a mere misunderstanding of our text, figures, and interpretation.

Morphology and internal structure

We do not imply, have not argued, and did not show in our figures that the Oxaya Block is a collapse feature in the sense of a chaotic landslide. In that respect, García and Hérail's arguments with reference to geometry, morphology, scale, volume, and internal structure are mute and comparison to the Lluta collapse or other paleo-landslides (e.g., Philip and Ritz, 1999) do not apply. Therefore, it is not an argument against the gravitational cause for rotation, that *'the Oxaya ignimbrites are coherent; they do not show a chaotic or disturbed structure with 'hummock' blocks and fragmented matrix, typical of gravitational slides and due to fragmentation during transport.'* (comment by García and Hérail). There are abundant examples for real landslides on the western (and eastern) slope of the Andes (e.g., Lluta Collapse, Naranjo *et al.*, 1993; Seyfried *et al.*, 1995; Hermanns and Strecker, 1999); they come in many sizes and shapes. It is obvious to us that the Oxaya Block is **not** a chaotic landslide and that this structure is larger by more than an order of magnitude than any of the landslides mentioned above.

Steepness of the western slope of the Andes in the Arica Bend area

It is not crucial to the discussion whether the slope of the Western Altiplano directly north and south of the Oxaya Block may be similarly steep, it certainly is. The high resolution morphological map compiled by the Cornell Andes group, however, does show that N and S of the Oxaya structure the distance between the Western Cordillera and the coast line does in fact increase (http://www.geo.cornell.edu/geology/cap/CAP_gen/CAP_topo.html). Therefore, the slope from the Western Cordillera to the coast is steeper in the Arica Bend area (Figs. 1b, c).

Shape and steepness of the 'Belén Escarpment'

The significance of the steep escarpment between the Oxaya structure and the crest of the Western Cordillera is in fact crucial for our interpretation. In relative terms, the Belén-Escarpment is in fact very steep. Outcrops of Oxaya Ignimbrites to the north and south of the Oxaya Block have not been differentially eroded in significant amounts. Retreating erosion, however, has smoothed the Belén Escarpment. Originally, steepness was not only larger in absolute terms but also relatively to the gently tilted Oxaya surface. We maintain our statement that the Belén Escarpment, which exposes some of Chile's oldest dated rocks (e.g., Wörner *et al.*, 2000b), represents an unusually steep portion of the Western Andean slope in northern Chile. It probably was even steeper prior to retreating erosion, and it should have genetic connection to the origin of the Oxaya Block, which it overtops.

Lateral and frontal boundaries of the Oxaya Block

We show in figure 1 that the lateral boundaries of the Oxaya Block are in fact transitions rather than boundaries. The overall shape and movement of the block, as seen in the 3-D satellite image is that of a large sag with accommodation zones to the surrounding regions. Therefore, we do not expect sharp sinistral (in the south) and dextral (in the north) offset boundaries, as demanded by García and Hérail.

INTERPRETATIONS OF THE OXAYA BLOCK

There are two fundamentally different models to explain the rotation of this giant block: García and Hérail's interpretation calls for a tectonic ramp structure bounded by two west-vergent thrusts (García *et al.*, 1999, and Comment by García and Hérail). Alternatively, the Oxaya Block is an antithetically rotated, giant gravitational block resulting from oversteepening of the western Andean slope in the Arica Bend area (Wörner *et al.*, 2000).

The sheer size of the Oxaya Block may turn out to be the main problem in accepting it as a large gravitational structure. However, large, gravitationally driven cover nappes are common wherever parts of the Earth's crust has been isostatically uplifted (e.g., Austroalpine and Penninic cover nappes upon and Helvetic cover nappes within the external zone of the Alps).

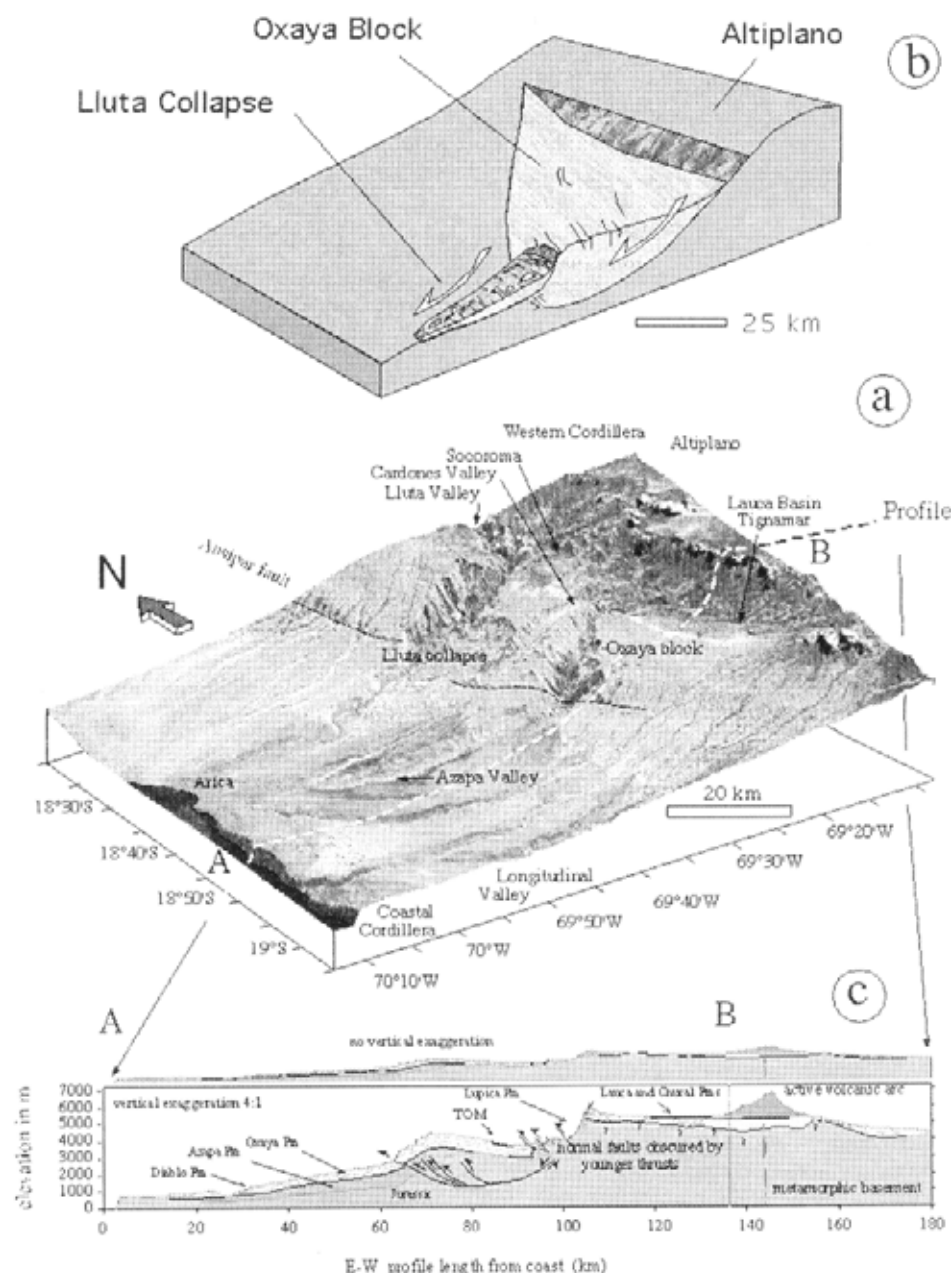


FIG. 1. **a**- 3-D satellite image showing the structure and morphology of the Oxaya Block and the steep Belén-Escarpment. Note that the lateral boundaries are very smooth and the north and south tips of the structure are actually at a higher elevation than the central parts. The simplified morphological model, **b**- depicts the block rotation and the Luta collapse at the steepened front; **c**- west-east cross section through the Oxaya Block along the topographical profile taken from the 1:250,000 topographical map. The profile shows the steep Belén-Escarpment.

In the following we justify why we do consider the ramp-thrust interpretation to be in conflict with some observations. Secondly, we present data that strongly favour our concept of gravitationally driven tectonics.

Problems with the ramp-and-thrust model

One problem to explain such structure by a stack of ramp thrusts is the fact, that this process would produce much steeper limbs than are actually observed (R. Allmendinger, oral communication, 2000). Rocks of similar competence, which deformed in a ramp-and-thrust fashion (e.g., Kley, 1996) indeed have limbs significantly steeper than observed here.

Another problem is the presumed short duration of the rotation (10.6–12 Ma) which is more consistent with a gravitational **event** rather than continuous tectonic movements related to the uplift of the Altiplano. Rapid displacement and rotation are further constrained by radiometric ages provided by Wörner *et al.* (2000) and the observation of a reversed drainage system on the Oxaya Block (Uhlig, 2000; Wörner *et al.*, in press).

Both Muñoz and Charrier (1996) and the Comment by García and Hérail imply that the reverse faults in question mainly caused the uplift of the western Altiplano. We see two problems with this interpretation:

- 1- the amount of uplift (1500 m) is not explained by the observed vertical throw along the faults (see below) and,
- 2- the morpho-tectonic structure of the Oxaya Block is restricted to the Arica bend area whereas uplift and a gently dipping western margin of the Altiplano is observed for many hundreds of kilometres alongstrike the western Andean slope. With reference to the slope of the Oxaya surface on both sides (north and south) of the Oxaya Block, the uplifted frontal part of the structure is geometrically balanced by the downward movement of the block. Thus, there is no net uplift related to the structure. While we do not dismiss westward thrusting along the western Andean margin, we observe that regional uplift is not confined to the existence of a ramp-and-thrust structure and that to the north and south of the Oxaya Block, the Western Cordillera is just as high.

Problems with our interpretation of a giant gravitational sag

A problem with our interpretation of the Oxaya Block as a gravitational structure was in fact accentuated by the Comment by García and Hérail in pointing out that where the Ausipar Thrust crops out (Lluta Valley), it is a concave-downward fault, shallowing to a subhorizontal thrust near the surface (García *et al.*, 1999). We do not observe the continuation at depth of the Ausipar Fault as a zone of decollement, but nor do we see the continuation as a blind thrust as postulated in García and Hérail's comment. The main problem with the interpretation of the Ausipar fault as the surface expression of a decollement is the fact, that it has a much smaller displacement than the offset between the crest of the Western Cordillera and the Oxaya Block (1,500 m, see above). This difference in movement between the east- and west-bounding fault zones would argue for internal deformation of the Oxaya Block (Fig. 1c). Moreover, the western limit of the Oxaya Block is close to the Ausipar reverse fault but it may not even be directly related. This is because the reverse offset of the Ausipar fault is larger in the N of the Oxaya Block (as exposed in the Lluta Valley, García *et al.*, 1996) compared to further south in the Azapa Valley.

We conclude from the discussion above that

- 1- Our model of a giant gravitational sag, as depicted in figure 1 was not fully explained in the paper by Wörner *et al.* (2000) and thus several important points were insufficiently documented and/or misinterpreted;
- 2- the ramp-and-thrust model is incompatible with some observations, and thus
- 3- the gravitational sag model is still a viable alternative.

We therefore reiterate our ...

Arguments in favour of a gravitational cause for Oxaya Block rotation

The Oxaya Block formed in the Arica bend area where the western Andean slope from coast to the crest of the Western Cordillera is by far the steepest.

The topography in this region produces a strong negative component in Gephart's (1994) antisymmetric residual in the overall Andean topographic symmetry. It is thus characterized by a topographic anomaly. While the present slope could be steeper than at the time when the Oxaya Block formed, we would argue that the particular location at the Arica Bend would have caused a relatively steep slope also in the geological past.

The short duration of movement (less than 2 Ma) for Oxaya Block rotation is consistent with a gravitational event rather than longer-lasting movements along reverse faults of a ramp and thrust structure.

With reference to the slope of the Oxaya surface to the north and south of the Oxaya Block, the uplifted frontal part of the structure is geometrically balanced by the downward movement of the block. Thus, there is no net uplift related to the structure as would be expected if the block were related to a tectonic ramp structure.

The Oxaya Block has its center at the altitude of the Cardones Valley. However, the Ausipar fault has a larger displacement N of the Lluta valley compared

to the area south of the Azapa valley. Thus, the Oxaya Block and the displacement of the Ausipar reverse fault are not directly related.

Our interpretation of a large tilted structure in an overall tectonic regime of convergence and regional tilting is also consistent with conclusions derived by Hartley *et al.* (2000) based on their tectonic and sedimentological analysis of the northern Chilean forearc. Isacks (1988); Lamb *et al.* (1997), and Lamb and Hoke (1997) also argued for regional west-ward tilting of the western Andean slope to explain the uplift of the western Altiplano. These models do not preclude but rather imply distributed reverse surface faulting. The mechanical model of Bailey (1998) also predicts gravitationally induced thrusting on oversteepened orogenic slope and ductile flow within the crust when rheologic properties of the crust are favourable. Rheological weakening due to heating by extensive magmatic activity is certainly attained in the Andean arc region.

One more observation in favour of a gravitational sag structure comes from off-shore geophysical investigations: the CINCA (1995) study showed the

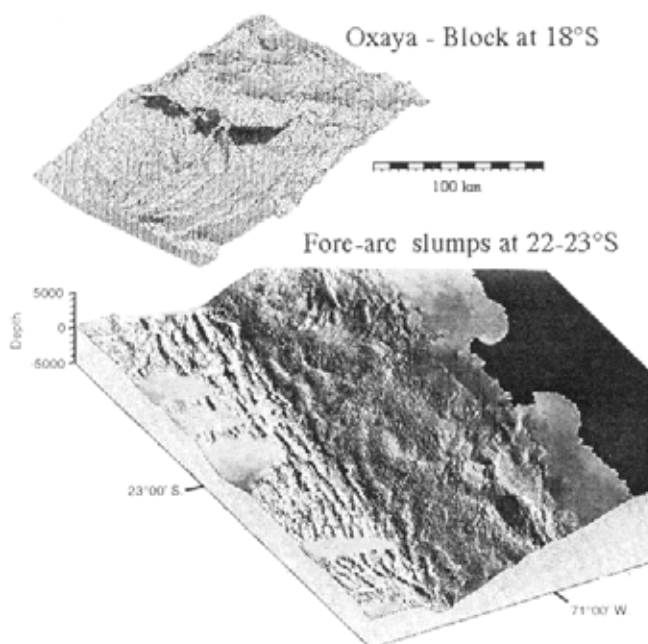


FIG. 2. Comparison of the digital elevation model of the fore-arc region near Antofagasta (modified after CINCA, 1995; see also von Huene *et al.*, 1999). Sizes and shapes of large rotated blocks are similar to the Oxaya Block and the fore-arc 'sags' below sea level. The latter have been identified by von Huene as antithetically rotated blocks due to oversteepening of the fore-arc region.

existence of large north-south oriented sagged blocks in the forearc (>30 km by 15 km in size, Fig. 2). These were interpreted to result from frontal (and basal) tectonic erosion of the outer forearc and subsequent antithetical gravitational rotation (von Huene *et al.*,

1999). In our opinion the Oxaya Block is an on-shore equivalent of these structures.

By weighing the above arguments, we still prefer the interpretation of the Oxaya Block as a large gravitational structure.

REFERENCES

- Allmendinger, R.W.; Jordan, T.E.; Kay, S.M.; Isacks, B.L. 1997. The Evolution of the Altiplano-Puna Plateau of the Central Andes. *Annual Reviews of Earth and Planetary Sciences*, Vol. 27, p. 139-174.
- Bailey, R.C. 1998. Thresholds for gravitationally induced thrusting by elevated topography over a ductile crust. Abstracts with Programs. *Geological Society of America*, Vol. 30, 297 p.
- CINCA. 1995. Crustal investigations off- and onshore Nazca/Central Andes. CINCA, Sonne-Fahrtbericht SO 104/3, Bundesanstalt für Geowissenschaften und Rohstoffe, p. 1-147. Hannover.
- Gaupp R.; Kött A.; Wörner, G. 1999. Paleoclimatic implications of Mio-Pliocene sedimentation in the high-altitude intra-arc Lauca Basin of northern Chile. *Palaeogeography, Palaeoclimatology, Palaeoecology*, Vol. 151, p. 79-100.
- García, M.; Hérial, G.; Charrier, R. 1999. Age and structure of the Oxaya Anticline: a major feature of the Miocene Compressive structures of northernmost Chile. Fourth ISAG Meeting, Göttingen, Germany. *Andean Geodynamics, IRD*, p. 249-252. Paris.
- García, M.; Hérial, G.; 2001. Comment on 'Geochronology (Ar-Ar, K-Ar and He-exposure ages) of Cenozoic magmatic rocks from northern Chile (18-22°S): implications for magmatism and tectonic evolution of the central Andes' of Wörner *et al.* 2000. *Revista geológica de Chile*, Vol. 28, No. 1p. 127-130, this issue.
- Gephart, J.W. 1994. Topography and subduction geometry in the Central Andes: Clue to mechanics of a noncollisional orogen. *Journal of Geophysical Research*, Vol. 99, p. 12279-12288.
- Gregory-Wodzicki, K.M. 2000. Uplift history of the Central and Northern Andes; a review. *Geological Society of America, Bulletin*, Vol. 112, p. 1091-1105.
- Hartley, A.J.; Geoffrey, M.; Chong, G.; Turner, S.; Kape, S.J.; Jolley, E. 2000. Development of a continental forearc; a Cenozoic example from the Central Andes, northern Chile. *Geology*, Vol. 28, p. 331-334.
- Hermanns, R.L.; Strecker, M.R. 1999. Structural and lithological controls on large Quaternary rock avalanches (sturzstroms) in arid northwest Argentina. *Geological Society of America, Bulletin*, Vol. 111, p. 934-948.
- Isacks, B.L. 1988. Uplift of the Central Andean Plateau and Bending of the Bolivian Orocline. *Journal of Geophysical Research*, Vol. 93, p. 3211-3231.
- Kley, J. 1996. Transition from basement-involved to thin-skinned thrusting in the Cordillera Oriental of southern Bolivia. *Tectonics*, Vol. 15, p. 763-775.
- Lamb, S.; Hoke, L. 1997. Origin of the high plateau in the Central Andes, Bolivia, South America. *Tectonics*, Vol. 16 p. 523-649.
- Lamb, S.; Hoke, L.; Kennan, L.; Dewey, J. 1997. Cenozoic evolution of the Central Andes in Bolivia and northern Chile. In *Orogeny through time* (Burg, J.P.; Ford, M.; editors). *Geological Society of London, Special Publication*, Vol. 121, p. 237-246.
- Muñoz, N.; Charrier, R. 1996. Uplift of the western border of the Altiplano on a west-vergent thrust system, Northern Chile. *Journal of South America Earth Sciences*, Vol. 9, p. 171-181.
- Naranjo J.A. 1993. Hallazgo de una gigantesca avalancha de detritos del Cenozoico superior en Oxaya, Región de Tarapacá. *Segundo Simposio Internacional de Estudios Altiplánicos, Actas*, p. 47-52. Arica, Chile.
- Philip, H.; Ritz, J.-F. 1999. Gigantic paleo-landslide associated with active faulting along the Bogd Fault (Goby-Altay, Mongolia). *Geology*, Vol. 27, p. 211-214.
- Salas, R.; Kast, R.F.; Montecinos, F.; Salas, I. 1966. Geología y recursos minerales del Departamento de Arica. Provincia de Tarapacá. *Instituto de Investigaciones Geológicas, Boletín*, Vol. 21, 114 p.
- Seyfried, H.; Wörner, G.; Uhlig, D.; Kohler, I.; Calvo, C. 2000. La geología y morfología de los Andes en el norte de Chile. *Universidad de Tarapacá, Chungará*, Vol. 30, p. 7-39. Arica, Chile.
- Seyfried, H.; Wörner, G.; Uhlig, D.; Kohler, I. 1995. Eine kleine Landschaftsgeschichte der Anden in Nordchile. Wechselwirkungen. *Jahrbuch der Universität Stuttgart*, p. 60-72.
- Uhlig, D. 1999. Die Westabdachung der Zentrallanden in den Provinzen Arica und Parinacota, Nordchile: Landschaftsentwicklung und Geologie. Dissertation. *University of Stuttgart, Profil*, Vol. 17, p. 167-244.
- von Huene, R.; Weinrebe, W.; Heeren, F. 1999. Subduction erosion along the North Chile margin. *Journal of Geodynamics*, Vol. 27, p. 345-358.
- Wörner, G.; Seyfried, H.; Kohler, I.; Uhlig, D. 1999. Uplift and erosion at the Western Andean Escarpment (WARP) in Northern Chile. In *International Symposium on Andean Geodynamics, No. 4, Extended Abstracts Volume*, p. 810-814. Göttingen.

- Wörner, G.; Hammerschmidt, K.; Henjes-Kunst, F.; Lezaun, J.; Wilke, H. 2000a. Geochronology (Ar-Ar, K-Ar and He-exposure ages) of Cenozoic magmatic rocks from northern Chile (18-22°S): Implications for magmatism and tectonic evolution of the central Andes. *Revista Geológica de Chile*, Vol. 27, No. 2, p. 205-240.
- Wörner, G.; Lezaun, J.; Beck, A.; Heber, V.; Lucassen, F.; Zinngrebe, E.; Rößling, R.; Wilke, H.G. 2000b. Precambrian and Early Paleozoic evolution of the Andean basement at Belén (N. Chile) and C. Uyarani (W. Bolivian Altiplano). *Journal of South American Earth Sciences*, Vol. 13 p. 717-737.
- Wörner, G.; Uhlig, D.; Kohler, I.; Seyfried, H. in press. Evolution of the West Andean Escarpment at 18°S (N. Chile) during the last 25 Ma: Uplift, erosion and collapse through time. *Tectonophysics*.